

# Robust identification and characterization of thin soil layers in cone penetration data by piecewise layer optimization

Jon Cooper, Eileen R. Martin, Kaleigh M. Yost, Alba Yerro, Russell A. Green

*Virginia Tech, Blacksburg, VA, USA*

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## Abstract

Cone penetration testing (CPT) is a preferred method for characterizing soil profiles for evaluating seismic liquefaction triggering potential. However, CPT has limitations in characterizing highly stratified profiles because the measured tip resistance ( $q_c$ ) of the cone penetrometer is influenced by the properties of the soils above and below the tip. This results in measured  $q_c$  values that appear “blurred” at sediment layer boundaries, inhibiting our ability to characterize thinly layered strata that are potentially liquefiable. Removing this “blur” has been previously posed as a continuous optimization problem, but in some cases this methodology has been less efficacious than desired. Thus, we propose a new approach to determine the corrected  $q_c$  values (i.e. values that would be measured in a stratum absent of thin-layer effects) from measured values. This new numerical optimization algorithm searches for soil profiles with a finite number of layers which can automatically be added or removed as needed. This algorithm is provided as open-source MATLAB software. It yields corrected  $q_c$  values when applied to computer-simulated and calibration chamber CPT data. We compare two versions of the new algorithm that numerically optimize different functions, one of which uses a logarithm to refine fine-scale details, but which requires longer calculation times to yield improved corrected  $q_c$  profiles.

*Keywords:* cone penetration test, data quality, inverse problems

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## 1. Introduction and Motivation

Cone penetration testing (CPT) is a preferred method to characterize soil profiles to evaluate seismic liquefaction triggering potential. The test consists of hydraulically pushing an instrumented cone-shaped penetrometer into the soil profile at a constant rate, with measurements typically taken every one to two centimeters as the cone advances. In its basic form, CPT sounding data include tip resistance ( $q_c$ ) and sleeve friction ( $f_s$ ) as a function of depth (Schmertmann, 1978). CPT  $q_c$  profiles are extensively used in geotechnical applications, in particular, they serve as a proxy for a soil’s ability to resist liquefaction triggering due to ground shaking (Shibata and Teparaksa, 1988).

CPT  $q_c$  profiles do not provide truly depth-specific measurements, because they are influenced by soil materials several cone diameters away from the cone tip (Ahmadi and Robertson, 2005). Consequently, if soil properties vary with depth, the measured  $q_c$  are “blurred” compared to the actual depth-specific or “true” corrected  $q_c$  (i.e., the  $q_c$  value that would be measured at that depth in a uniform profile, absent of boundary or thin-layer effects). This “blurring” is asymmetrical, with soil below the cone tip affecting  $q_c$  more strongly than soil above, as illustrated in Figure 1 (Boulanger and DeJong, 2018). This figure shows a multi-layer “true”  $q_c$  profile along with the computationally simulated data we would generate by convolving the profile with an asymmetrical function. Notice how the peaks of the simulated data are shifted up relative to the true high  $q_c$  layers and how the true  $q_c$  in the thin layers is obscured. Because soil profiles are typically stratified, the location of the interfaces between layers and the true  $q_c$  of layers can become difficult to precisely identify from the measured  $q_c$ , even in relatively simple profiles. These phenomena are typically referred to as thin-layer, transition-zone, or multiple thin-layer effects, as discussed in Yost et al. (2021b). Herein, we will generally refer to all of these effects as “thin-layer effects”. In this context, a soil “layer” or “stratum” is a depth increment in the profile over which the soil has relatively uniform geotechnical engineering properties (e.g., soil type and  $q_c$ ). Furthermore, a “thin” layer or stratum is one that is too thin for the measured  $q_c$  to fully develop or reach values that would be measured in the stratum absent of thin-layer effects (i.e. true corrected  $q_c$  profile). This required thickness will vary as a function of soil stiffness, but typically 10 to 30 cone diameters is required (Ahmadi and Robertson, 2005), and thus a stratum thinner than this would be a “thin layer.”

The majority of studies on thin-layer effects on measured CPT  $q_c$  data

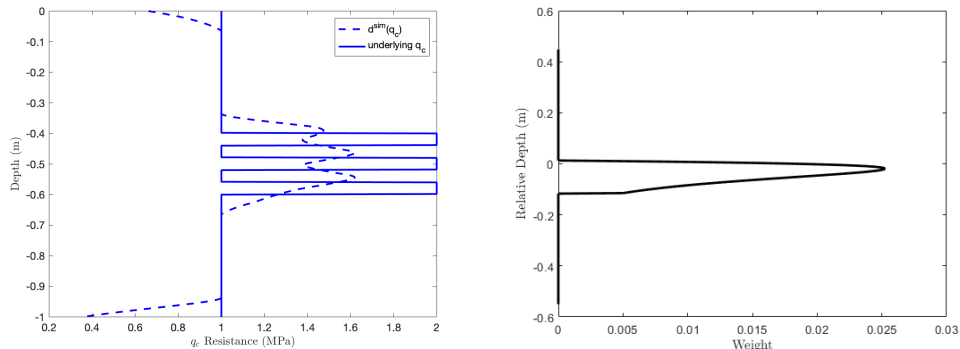


Figure 1: The  $q_c^{true}$  profile of a layered soil (left, solid) and its predicted blurred observation,  $\tilde{q}_c^{sim}$  (left, dashed), generated by convolving the profile with a point spread function that is a truncated chi-squared distribution (right).

38 have focused on developing corrections to apply to measured  $q_c$  for specific  
 39 layering sequences and geometries, as outlined in Boulanger and DeJong  
 40 (2018). Some past efforts to correct the measured  $q_c$  for thin-layer effects in  
 41 profiles consisting of a stiff layer embedded in a soft layer (e.g. sand and clay  
 42 layers, respectively) assumed that the effect of the softer layer will be greater  
 43 at the boundaries and lesser towards the middle of the stiff layer. This results  
 44 in a V-shaped deblurring correction function to correct the CPT-measured  
 45  $q_c$  in the stiff layer so the corrected values more closely represent the true  
 46  $q_c$  that would be measured in the absence of thin-layer effects (Youd et al.,  
 47 2001; de Greef and Lengkeek, 2018).

48 In comparison to corrections for specific layering scenarios, Boulanger  
 49 and DeJong (2018) propose a potentially more flexible inverse problem ap-  
 50 proach that can be applied to measured  $q_c$  from a CPT performed in a profile  
 51 containing any number of layers. However, the procedure is less efficacious  
 52 than desired in some scenarios where layer thicknesses are less than 40 mm,  
 53 even if there is significant contrast between the strengths and stiffnesses (i.e.,  
 54 the true  $q_c$ ) of the thin layers and surrounding soil (Yost et al., 2021a).  
 55 The inability to cost-effectively determine the true  $q_c$  from the measured  $q_c$   
 56 may be contributing to widespread over-prediction of the liquefaction haz-  
 57 ard in highly interlayered soil deposits (e.g., as observed in Christchurch,  
 58 New Zealand (Maurer et al., 2014, 2015) and the Hawk’s Bay region of New  
 59 Zealand (El Korthawi et al., 2019)). In this regard, a quantitative comparison  
 60 of previous correction procedures is presented in Yost et al. (2021a).

61 The main objectives of this work are (i) to pose a new inverse problem  
62 to estimate true  $q_c$  from measured (or “blurred”)  $q_c$  in highly stratified soil  
63 profiles, and (ii) to propose a new numerical optimization algorithm for ef-  
64 ficiently correcting  $q_c$  in highly stratified soil profiles for thin-layer effects.  
65 This paper is organized as follows: In Section 2, we review the background  
66 of prior efforts to adjust or correct CPT data for thin-layer effects, including  
67 efforts to pose this correction as an inverse problem to be solved via numer-  
68 ical optimization. Further, we provide relevant background information on  
69 numerical optimization techniques. In Section 3, we propose a new approach  
70 to correct CPT data by removing the thin-layer effects via an inverse prob-  
71 lem, posed in two different ways. We describe a new algorithm to solve both  
72 proposed versions of this inverse problem that incorporates global numerical  
73 optimization techniques, routines to automatically generate a reasonable ini-  
74 tial guess, to adjust the number of layers, and to computationally simulate  
75 the blurring process. In Section 4, we show the results of this new algorithm  
76 for one version of the inverse problem applied to CPT data from calibration  
77 chamber tests, and the improvements typically achieved over simpler com-  
78 putational methods to automatically correct thin-layer effects in measured  
79  $q_c$  profiles. In Section 5, we compare the results of the new algorithm for  
80 both formulations of the inverse problem, yielding a more precise, but more  
81 computationally expensive data correction method. In Section 6, we discuss  
82 some possible future extensions of this method, and in Section 7, we discuss  
83 limitations and benefits of the new algorithm.

## 84 2. Background

85 In computational science and engineering, the terms “forward problem”  
86 and “inverse problem” are often used to describe the problems we solve when  
87 different parts of a system are unknown. When we assume we know the sub-  
88 surface soil characteristics and stratigraphy of a soil profile (i.e., we assume  
89 an unblurred  $q_c$  profile), and then perform a computational simulation to pre-  
90 dict the response of the soil to an action (e.g., simulating a CPT in a “known”  
91 (unblurred)  $q_c$  profile to compute a simulated blurred tip resistance profile,  
92  $\tilde{q}_c^{sim}$ , comparable to a measured  $q_c$  profile for a real CPT), we are solving the  
93 “forward problem.” In contrast, we are solving the “inverse problem” when  
94 we only know the measured tip resistance profile,  $\tilde{q}_c^{meas}$ , but need to infer  
95 the true soil characteristics and stratigraphy (i.e.,  $q_c^{true}$ ) that would lead to  
96 computationally simulated data,  $\tilde{q}_c^{sim}$ , that most closely match the measured

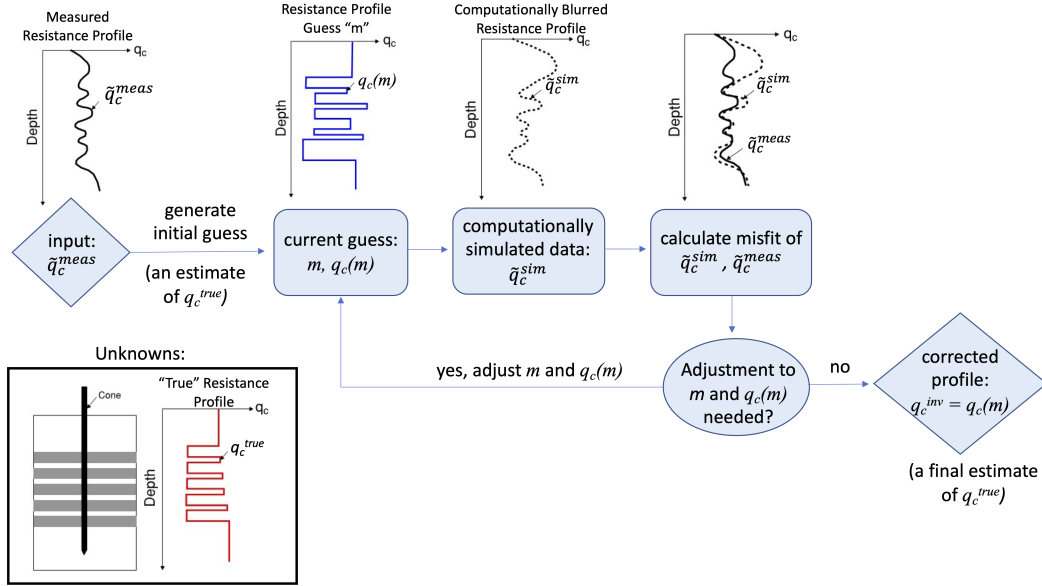


Figure 2: Diagram of the inverse problem approach to remove thin-layer effects.

97 tip resistance  $\tilde{q}_c^{meas}$ . In this inverse problem scenario, we begin with some  
 98 initial guess at  $q_c$  and iteratively update the current guess to incrementally  
 99 improve the match between  $\tilde{q}_c^{meas}$  and  $\tilde{q}_c^{sim}$  until reaching "the best" guess  
 100 (i.e., the solution to the inverse problem), denoted by  $q_c^{inv}$ .

101 In this regard, the inverse problem is posed as an optimization problem  
 102 (i.e., were the minimum difference between  $\tilde{q}_c^{meas}$  and  $\tilde{q}_c^{sim}$  is targeted) and  
 103 solved using a variety of numerical optimization techniques. All numerical  
 104 optimization techniques require calculating  $\tilde{q}_c^{sim}$  for every  $q_c$  guess, so we  
 105 solve the forward problem many times to solve the inverse problem once.  
 106 We can predict  $\tilde{q}_c^{sim}$  for any  $q_c$  either through i) numerical simulation of the  
 107 soil being displaced by the cone penetrometer (e.g., with numerical methods  
 108 such as the material point method (Yost et al., 2021b)), or ii) applying a  
 109 simplified blurring function to  $q_c$ . This workflow, beginning with  $\tilde{q}_c^{meas}$ , iter-  
 110 atively updating the proposed  $q_c$  and its corresponding  $\tilde{q}_c^{sim}$ , and ultimately  
 111 outputting the corrected  $q_c^{inv}$ , is diagrammed in Figure 2.

112 Approaches to solve inverse problems have previously been applied to  
 113 other geotechnical engineering challenges. Notably, the multichannel analysis  
 114 of surface waves (MASW) technique for seismic imaging of the near surface is  
 115 an inverse problem (see Socco and Strobbia (2004) for details). However, the

116 development of thin-layer corrections for CPT data collected in interlayered  
 117 soil profiles has only recently been posed as an inverse problem by Boulanger  
 118 and DeJong (2018). In the following, we present some background on these  
 119 methods.

## 120 2.1. Prior Methods and Limitations

121 Regardless of the method to calculate  $\tilde{q}_c^{sim}$ , this inverse problem can be  
 122 generically posed as finding the assumed  $q_c$  that minimizes a function known  
 123 as a *misfit function*:

$$q_c^{inv} := \arg \min_{q_c} \|\tilde{q}_c^{meas} - \tilde{q}_c^{sim}(q_c)\| \quad (1)$$

124 where  $\tilde{q}_c^{meas}$ ,  $q_c$ , and  $\tilde{q}_c^{sim}$  are all vectors with as many entries as there are  
 125 depths of interest. The tilde indicates profiles that are “blurry” while the  
 126 lack of a tilde indicates profiles that are constrained to be layered profiles  
 127 with sharp transitions. The misfit function (Eq. 1) measures the difference  
 128 between the actual measured profile and the simulated measured profile for  
 129 any  $q_c$  guess. An engineers’ physical understanding of the influence of thin-  
 130 layer effects on measured CPT data can be incorporated by using a physically  
 131 realistic computational simulation process (i.e., “blurring function”) to map  
 132 the current guess at the resistance profile,  $q_c$ , to  $\tilde{q}_c^{sim}$ . However, inverse  
 133 problems are often ill-posed and data include noise, so it is possible that  
 134 a small change in  $\tilde{q}_c^{meas}$  could allow for significantly different  $q_c^{inv}$  profiles  
 135 that both yield  $\tilde{q}_c^{sim}$  equally close to  $\tilde{q}_c^{meas}$ . By modifying the way that we  
 136 discretize and mathematically represent the soil profile, changing the form  
 137 of the misfit function, adding physically realistic restrictions on  $q_c$  that are  
 138 considered, or using different numerical optimization algorithms to iteratively  
 139 improve the corrected  $q_c$  guesses, we may be able to improve the efficacy of  
 140 our solution.

141 This inverse problem approach to correcting  $q_c$  for thin-layer effects was  
 142 first used by Boulanger and DeJong (2018). Their key insight was repre-  
 143 senting thin-layer effects on  $\tilde{q}_c^{meas}$  at a particular depth as a simple blurring  
 144 filter,  $w_c(z)$ , applied to the true tip resistance profile  $q_c^{true}$ . They inherently  
 145 assumed that the coefficients of the blurring function may vary with depth.

$$\tilde{q}_c^{sim}(q_c) := q_c * w_c(z) \quad (2)$$

146 where  $*$  represents convolution. It is assumed that  $w_c$  is the discretization  
 147 of a continuous function that represents the influence of soil above and be-

148 low the cone tip on  $\tilde{q}_c^{meas}$  at a particular depth. An example of a  $w_c$  func-  
 149 tion that is a scaled and truncated chi-squared distribution is shown along  
 150 with the layered soil  $q_c$  and  $\tilde{q}_c^{sim}$  in Figure 1. This differs from the  $w_c$  used  
 151 in Boulanger and DeJong (2018), but our numerical tests suggest this  $w_c$   
 152 more closely matches calibration chamber data. The numerical optimization  
 153 method proposed in Boulanger and DeJong (2018) uses a common iterative  
 154 splitting optimization technique, and smooths the results to keep them from  
 155 becoming unstable. However, when applied to laboratory calibration cham-  
 156 ber test data, we found that this method may still be unstable (i.e., it did not  
 157 yield corrected  $q_c$  profiles that matched the stratigraphy of known soil profiles  
 158 with thin layers) (Yost et al., 2021a). We have explored a variety of ways  
 159 to pose this inverse problem as different optimization problems, methods to  
 160 discretely represent the soil profiles, and numerical optimization techniques  
 161 including both gradient-based methods and global optimization techniques.  
 162 In this paper we i) propose two new representations of this inverse problem,  
 163 ii) detail a new robust numerical optimization algorithm for each represen-  
 164 tation to find the best guess for the resistance profile,  $q_c^{inv}$ , iii) present the  
 165 tradeoffs in accuracy and computational cost of these algorithms, and iv)  
 166 provide open source MATLAB code for all algorithms and test cases.

### 167 3. New Method

168 We generally assume each soil layer is homogeneous and refer to the  
 169 “corrected tip resistance” in this homogeneous layer as the tip resistance  
 170 that would be measured in an entirely uniform profile of the same material  
 171 (perhaps with some level of noise). Accordingly, our new method describes  
 172 any proposed  $q_c$  profile as a piecewise constant function. Assuming there are  
 173  $N$  soil layers, each layer in the piecewise constant function is represented by  
 174 just two degrees of freedom, rather than having as many degrees of freedom  
 175 as number of depths where CPT data were collected. The two variables  
 176 associated with each layer would be i) its depth, and ii) its characteristic tip  
 177 resistance if the uniform material in the layer were measured without any  
 178 thin-layer effects. This representation results in far fewer degrees of freedom  
 179 compared to the formulation of Boulanger and DeJong (2018), and improves  
 180 computational speed of the method. However, even simple soil profiles can  
 181 have several dozen degrees of freedom (i.e., two degrees of freedom per layer),  
 182 requiring a global optimization method that balances efficiency and precision.

183 We define the inverse problem to seek the  $q_c$  that results in  $\tilde{q}_c^{sim}$  that

184 most closely matches  $\tilde{q}_c^{meas}$ . We restrict  $q_c$  to be a piecewise constant function  
 185 defined by  $N$  layer depths paired with  $N$   $q_c$  values. Therefore, each piecewise  
 186 proposed  $q_c$  profile is described by a material property vector,  $m$ , that has  $2N$   
 187 components. For any assumed  $m$  we can reconstruct the corresponding  $q_c$  at  
 188 each depth where CPT data were measured by simply extracting the value  
 189 of the piecewise function described by  $m$  at every depth of interest. The  $q_c$   
 190 profile resulting from this reconstruction process is denoted by  $q_c(m)$ . In what  
 191 we refer to as the *new algorithm with the standard misfit function*, we quantify  
 192 how closely the measured and simulated profiles match by calculating  $q_c(m)$ ,  
 193 calculating  $\tilde{q}_c^{sim}$  from  $q_c(m)$ , then calculating the norm (the Euclidean or 2-  
 194 norm) of the error between  $\tilde{q}_c^{sim}$  and  $\tilde{q}_c^{meas}$ , scaled to be between 0 and 1. For  
 195 reference, a good fit would have a score of less than 0.01 (i.e., less than 1%  
 196 error). By convention, we consider the depth of the layer to be the depth  
 197 to the top of the layer, and we also force the depth to the first layer to be  
 198 zero. Written as an equation, this new algorithm with the standard misfit  
 199 function optimizes:

$$m^{inv} = \arg \min_{m \in \mathbb{R}^{2N}} \|\tilde{q}_c^{meas} - \tilde{q}_c^{sim}(q_c(m))\|_2. \quad (3)$$

200 However, this is not the only way to quantify the misfit between  $\tilde{q}_c^{meas}$  and  
 201  $\tilde{q}_c^{sim}$ . Some inverse problems that have data or material parameters that in-  
 202 clude both small-scale and large-scale values benefit from quantifying misfits  
 203 with a logarithm applied. Thus we also propose the *new algorithm with the*  
 204 *log misfit function*, which optimizes:

$$m^{inv} = \arg \min_{m \in \mathbb{R}^{2N}} \log(\|\tilde{q}_c^{meas} - \tilde{q}_c^{sim}(q_c(m))\|_2). \quad (4)$$

205 Computationally optimizing either form of the misfit function based on mea-  
 206 sured and simulated data allows us to assess  $q_c$  guesses without direct knowl-  
 207 edge of  $q_c^{true}$ , even when additional site characterization data are unavailable.  
 208 The best assumed  $q_c$  profile we reconstruct,  $q_c^{inv} = q_c(m^{inv})$ , is likely to be  
 209 close to the  $q_c^{true}$  profile with thin-layer effects removed, but practical numer-  
 210 ical optimization algorithms may yield different answers depending on the  
 211 choice of the misfit function.

212 In addition to designing the optimization problem, one must select a  
 213 numerical optimization algorithm to iteratively update the  $q_c$  guesses. We  
 214 chose the Particle Swarm Optimization (PSO) algorithm. PSO finds minima  
 215 of the selected misfit function starting with many randomly generated trial



216  $m$  values (i.e., the “particles”), each following its own path of new updated  
217 guesses of the  $q_c$  profile (i.e. guesses for  $m$  with a corresponding  $q_c(m)$ ). Each  
218 particle explores the space of possible  $m$  vectors based on its most recent  $m$   
219 guess, the best  $m$  guess it has tested, and the best  $m$  guess previously tested  
220 among all the particles. In this way, PSO does indeed have particles that  
221 swarm around local and global minima.

222 Since PSO does very well when the global minimum is surrounded by  
223 local minima or has a wide basin of attraction (i.e. a large region around the  
224 global minimum with no other local minima), and small adjustments to either  
225 the layer depths or resistances should only marginally affect  $\tilde{q}_c^{meas} - \tilde{q}_c^{sim}$ , we  
226 believe PSO is a practical choice. When  $\tilde{q}_c^{meas}$  does not suggest a piecewise  
227 constant layer resistance profile (e.g., when there is a gradient in the  $q_c$   
228 profile), we can still approximate the result well by adding several additional  
229 layers, each with constant  $q_c$ . In this case, there may be multiple local minima  
230 surrounding the global minimum, and so again PSO should be quite effective.  
231 The only drawback to PSO is that it can have low accuracy, i.e., different  
232 initial particle guesses may yield quite different  $q_c^{inv}$  values even if the average  
233 over all particles’  $q_c$  guesses is at the global minimum.

234 The pseudocode for this new algorithm is described in Algorithm 1 be-  
235 low. The following subsections step through the process to automatically  
236 compute good initial  $m$  guesses, followed by two methods used in tandem  
237 with PSO for adjusting the number of layers and re-fitting the  $q_c$  profile  
238 guess automatically. The pseudocode assumes that the user has already se-  
239 lected whether they wish to use the standard or log misfit function. Further,  
240 while the pseudocode is written assuming use of the recommended initial-  
241 ization methods outlined in Section 3.1, an engineer applying this algorithm  
242 may substitute their own initial guess of  $q_c$  based on their knowledge of local  
243 soils and geology.

### 244 3.1. Calculating Reasonable Initial $m$ Guess

245 In order to combat the accuracy limitations of PSO, a standard tech-  
246 nique in optimization is to focus on developing a good initial  $m$  guess (and  
247 its corresponding  $q_c(m)$ , referred to here as the initial  $q_c$  guess) which can  
248 then be further refined by PSO. Many PSO implementations allow for the  
249 specification of the initial  $m$  guess. Even if a single particle starts at a good  
250 initial  $m$  guess, by the nature of PSO, the other particles will quickly swarm  
251 the location and discover the global minimum. Here we propose a novel tech-  
252 nique to automatically generate a good initial guess, which is specific to the

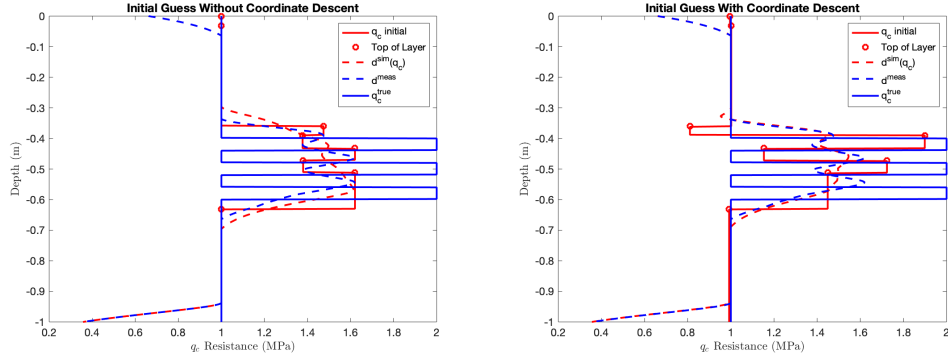


Figure 3: An initial guess for  $q_c$  (left, solid red) was automatically generated from derivative sign changes of the measured  $\tilde{q}_c^{meas}$  resistance profile (dashed blue). Its computationally simulated  $\tilde{q}_c^{sim}$  data is shown (left, dashed red). Coordinate descent optimization starting from that initial guess of  $q_c$  yielded a new initial guess (right, solid red) with an improved predicted blurred profile (right, dashed red).

253 thin soil layer problem.

254 The first step in constructing a good initial  $m$  guess (and its correspond-  
 255 ing  $q_c(m)$ ) is to automatically calculate an approximation of  $N$ , the number  
 256 of layers, and the depths of each layer. This can be done by looking at the  
 257 locations where the derivative of  $\tilde{q}_c^{meas}$  changes signs. This will not capture  
 258 features of  $q_c^{true}$  such as gradually increasing/decreasing resistances, but ef-  
 259 fectively this should identify most locations where there is a transition either  
 260 from a layer having a low  $q_c^{true}$  value to a high  $q_c^{true}$  value or vice-versa. The  
 261  $q_c$  values of the  $N$  layers can simply be initialized as being equal to the  
 262 measured resistances at a subset of the depths where  $\tilde{q}_c^{meas}$  was measured.

263 Although this initialization works much better than a random initializa-  
 264 tion, there is a chance that the asymmetry of the influence of the soil above  
 265 and below the cone tip on  $\tilde{q}_c^{meas}$ , and the number of thin layers might result in  
 266 an initial  $m$  and  $q_c(m)$  guess that are offset in depth from the true resistance  
 267 profile, as seen in Figure 3. To fix this, the new proposed algorithm applies  
 268 a simple coordinate descent optimization algorithm with the selected misfit  
 269 function to improve the initial  $m$  guess further at low computational cost.  
 270 Coordinate descent is a common numerical optimization technique (Wright,  
 271 2015). Applying coordinate descent optimization here helps to correct  $m$   
 272 when the locations of multiple layers are shifted from the true layer depths,  
 273 and improves the estimated  $q_c$  values slightly.

274 The result of these steps is a reasonable initial guess for  $q_c$  which, in rare  
275 cases, might already be optimal. However, coordinate descent optimization  
276 is typically unable to refine the details of  $q_c^{inv}$ , so coordinate descent is only  
277 used for a quick update to the initial guess followed by application of PSO  
278 and computational procedures to add/remove layers for further improvement.  
279 This is because PSO explores many more minima than just the single local  
280 minimum that coordinate descent finds. By using a good initial guess, PSO  
281 takes far less time to converge than with a random initial guess. These steps  
282 for constructing a good initial  $m$  guess (and its corresponding initial  $q_c$  guess)  
283 have no way of adjusting the number of layers in the profile, usually resulting  
284 in initial guesses that are too simplistic, so our proposed new algorithm  
285 includes computational procedures to automatically add and remove layers.

### 286 3.2. *Leave-One-Out (LOO)*

287 We propose a new computational procedure to improve a  $q_c$  guess by  
288 removing any layers that would help reduce the misfit function (either the  
289 standard or log misfit, depending on the user's choice) up to some pre-defined  
290 tolerance, referred to as the Leave-One-Out (LOO) procedure. To accomplish  
291 this, LOO computes what the misfit would be if the  $i^{th}$  layer were removed  
292 from the profile for each  $i = 1, 2, \dots, N$ , then removes whichever layer in-  
293 creases the misfit the least, up to the tolerance. This process is repeated until  
294 the removal of any single layer increases the misfit beyond the tolerance, or  
295 when the model contains only one layer. Algorithm 2 details the pseudocode  
296 of the LOO procedure.

297 LOO was designed to remove insignificant, if not detrimental, layers from  
298 any  $q_c$  guess that are not physically realistic and contribute to unnecessary  
299 additional degrees of freedom (which increase the run-time of) PSO. The  
300 provided software includes several options to set the tolerance automatically,  
301 most of which only rely on the misfit of the initial  $q_c$  guess and do not change  
302 between iterations.

### 303 3.3. *Add-One-In (AOI)*

304 In addition to removing unnecessary layers, we also developed a new  
305 computational procedure to automatically add missing layers, referred to  
306 as Add-One-In (AOI). AOI adds new layers between existing layers if the  
307 addition of that layer would reduce the misfit function at the proposed new  
308 profile, until the addition of another layer does not sufficiently decrease the  
309 misfit. This is a necessary condition to stop adding layers, since adding a

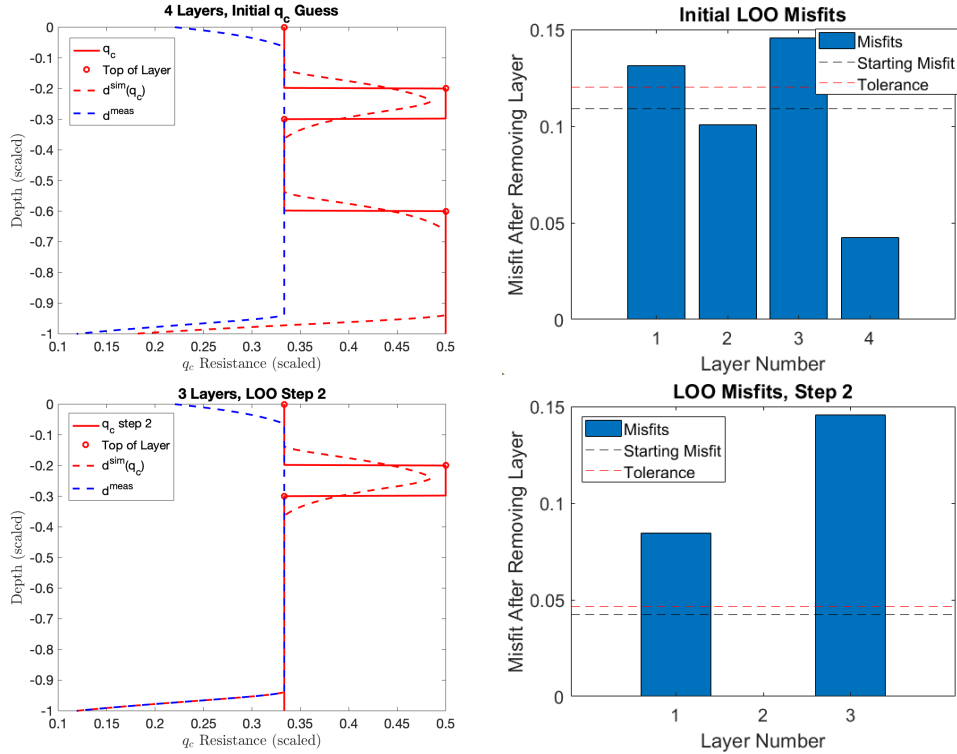


Figure 4: The first (top) and second (bottom) steps of the Leave-One-Out (LOO) process are demonstrated. At each step  $\tilde{q}_c^{sim}$  (red dashed) is calculated from the current guess at  $q_c$  (red solid), and compared to  $\tilde{q}_c^{meas}$  (blue dashed). The predicted misfit if each layer were to be removed is calculated (right). The first step has layers that can be removed to lower the misfit, but no layers are below the tolerance for removal in the second step.

310 layer is always guaranteed to at least keep the misfit the same, if not decrease  
 311 it, which potentially allows the number of layers in the profile to grow to  
 312 infinity without the stopping criteria. AOI accomplishes this by adding in a  
 313 layer between every two consecutive pairs of layers, improving these layers'  $q_c$   
 314 values and thicknesses using PSO (a quick, 2-variable optimization for each  
 315 layer), and computing which proposed additional layer decreases the misfit  
 316 the most. Algorithm 3 details the pseudocode of the AOI procedure. AOI was  
 317 designed to populate regions of high data misfit with more layers, assuming  
 318 the next full application of PSO will be able to adjust these new layers  
 319 appropriately. AOI struggles to add layers where multiple layers having a  
 320 mix of high and low  $q_c$  values are missing. Unlike LOO, AOI tolerances must  
 321 be updated each iteration to account for the potentially rapidly decreasing  
 322 misfit.

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**Algorithm 1** Thin-Layer Correction Optimization

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**Require:** measured data  $\tilde{q}_c^{meas}$ , and function to simulate data  $blurfcn$   
 $ndata \leftarrow \text{normalize}(\tilde{q}_c^{meas})$   
 create  $misfitfcn$ , a misfit function based on  $ndata$  and  $blurfcn$   
 initialize  $m$  based on where  $\text{deriv}(ndata)$  changes signs  
 $m \leftarrow \text{coordinateDescent}(m, misfitfcn)$   
**while**  $\text{length}(m) \neq \text{length}(m_O)$  **or**  $\|m - m_O\| > \varepsilon$  **or** first iteration **do**  
    $m_O \leftarrow m$   
    $m \leftarrow \text{LOO}(m, misfitfcn)$   
    $m \leftarrow \text{PSO}(m, misfitfcn)$   
    $m \leftarrow \text{AOI}(m, misfitfcn)$   
    $m \leftarrow \text{PSO}(m, misfitfcn)$   
**end while**  
 $m \leftarrow \text{LOO}(m, misfitfcn)$   
 check for potential uncertainties, state warnings  
 $m^{inv} \leftarrow \text{rescale } m \text{ to remove normalization}$   
 $q_c^{inv} \leftarrow \text{reconstruct depth profile } q_c(m^{inv})$   
**return**  $q_c^{inv}$

---

323 *3.4. Convolutional Blurring Procedure*

324 A user of the proposed new algorithm can use any method to calculate  
 325  $\tilde{q}_c^{sim}$  that they prefer. For the purposes of this study, we defined an arti-  
 326 ficial computational blurring function that, when applied to idealized  $q_c^{true}$ ,

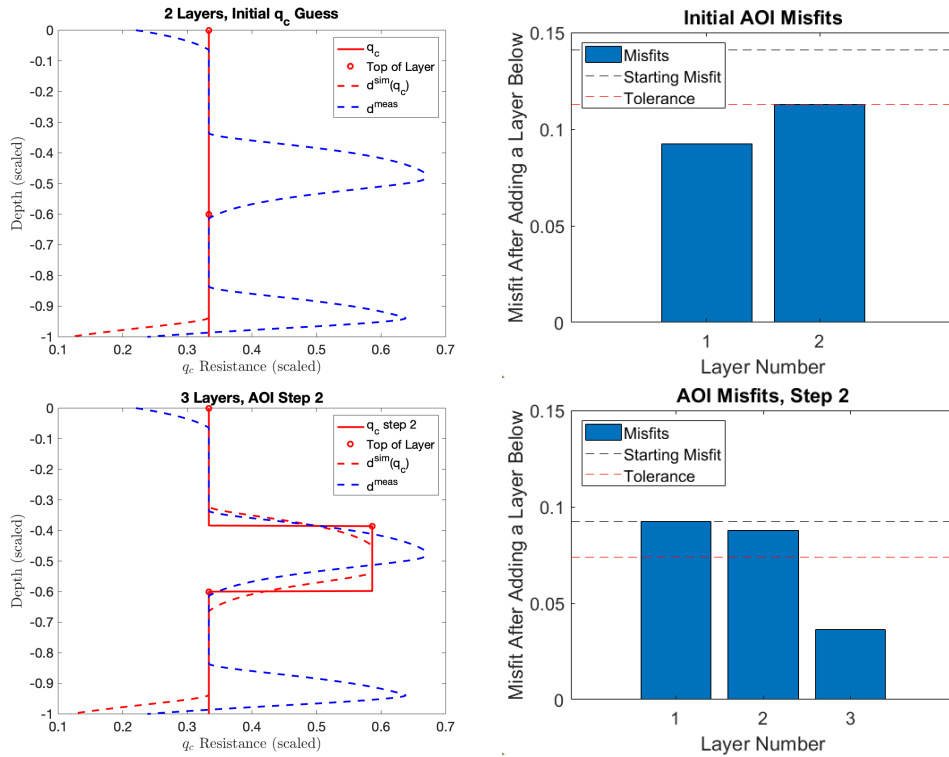


Figure 5: The first (top) and second (bottom) steps of the Add-One-In (AOI) process are demonstrated. At each step the predicted blurred data (red dashed) is calculated from the current guess at the resistance profile (red solid) and compared to the measured resistance profile (blue dashed). The predicted misfit assuming additional layers is calculated (right).

---

**Algorithm 2** Leave-One-Out (LOO)

---

**Require:** profile guess  $m$ , function to evaluate misfit  $misfitfcn$ , tolerance  $TOL$   
 $misfits \leftarrow \text{zeros}(N,1)$   
**while** true **do**  
  **if**  $N == 1$  **or**  $\text{all}(misfits > TOL)$  **then**  
    break  
  **end if**  
   $i \leftarrow$  index of minimum entry of  $misfits$   
   $m \leftarrow \text{removeLayer}(m, i)$   
   $N \leftarrow N - 1$   
  **for**  $i = 1 : N$  **do**  
     $temp \leftarrow \text{removeLayer}(m, i)$   
     $misfits(i) \leftarrow misfitfcn(temp)$   
  **end for**  
**end while**  
**return**  $m$

---

327 replicates the blurring caused by the thin-layer effects. That is, it compu-  
328 tationally simulates the  $\tilde{q}_c^{sim}$  from  $q_c^{true}$ , and the resulting values should be  
329 close to the actual  $\tilde{q}_c^{meas}$ . Similar to Boulanger and DeJong (2018) we chose  
330 to use a blurring function defined by convolving the true resistance profile  
331 with a *point spread function*  $p(z)$ :

$$\tilde{q}_c^{sim}(z) := (\tilde{q}_c^{sim}(q_c))(z) = \int_{-\infty}^{\infty} q_c(\Delta z) p(z - \Delta z) d\Delta z, \quad (5)$$

332 where  $\int_{-\infty}^{\infty} p(z) dz = 1$  and  $p(z) \geq 0$  for all  $z$ . In practice, this integral is only  
333 calculated on a finite interval. This blurring function was chosen because  
334 it is simple to implement and only requires the use of a matrix convolution  
335 function (“conv” in MATLAB). This blurring function results in a  $\tilde{q}_c^{sim}$  value  
336 at each depth that is a weighted combinations of the surrounding soil layers’  
337  $q_c$  values. Although this method can very quickly compute  $\tilde{q}_c^{sim}$  for any  $q_c$   
338 guess, it is a simplification of the true physics. For example, in regions with  
339 alternating thin layers of stiff and soft soils,  $q_c^{meas}$  in the layered zone is much  
340 closer to the lower of the two true resistances throughout the region (i.e., it  
341 is not a simple averaging process), as we see in Figure 6. Simplified physics  
342 models suggest that the layers below the tip of the cone affect the resistance

---

**Algorithm 3** Add-One-In (AOI)

---

**Require:** profile guess  $m$ , function to evaluate misfit  $misfitfcn$

```
while true do
  recompute  $TOL$ 
  if  $misfitfcn(m) \leq TOL$  then
    break
  end if
   $misfits \leftarrow \text{zeros}(N - 1, 1)$ 
  for  $i = 1 : N - 1$  do
    initialize  $extraLayer$  between layer  $i$  and layer  $i + 1$ 
     $temp \leftarrow \text{addLayer}(m, extraLayer, i)$ 
     $misfits(i) \leftarrow misfitfcn(temp)$ 
  end for
  if  $\text{all}(misfits) > TOL$  then
    break
  end if
   $i \leftarrow \text{index of minimum entry of } misfits$ 
   $m \leftarrow \text{addLayer}(m, extraLayer, i)$ 
end while
return  $m$ 
```

---



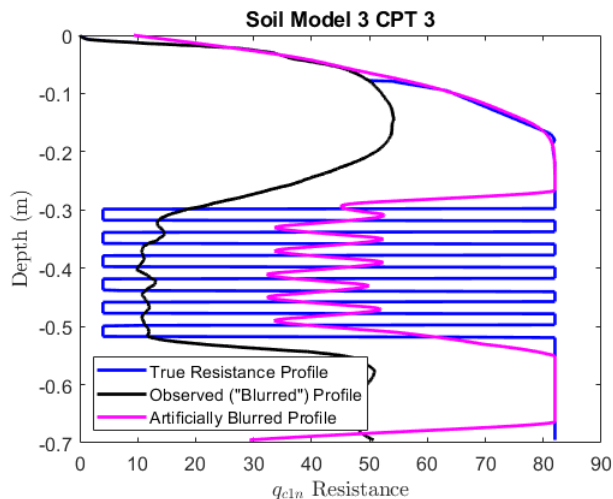


Figure 6: A CPT tip resistance profile from Soil Model 3 CPT 3 of the De Lange (2018) report, Section A. Here, the observed resistance profile  $\tilde{q}_c^{meas}$  (black) is being contrasted with  $\tilde{q}_c^{sim}$  (magenta) that is calculated from the true resistance profile  $q_c$  (blue) by simple convolution.

343 more than the layers above the cone tip (Boulangier and DeJong, 2018), so  
 344 in our implementation, we performed the convolution in Equation 5, with a  
 345 point spread function  $p(z)$  derived from the Chi squared probability density  
 346 function as the artificial blurring function for its smoothness and asymmetry,  
 347 pictured in Figure 1. Our computational experiments to find the optimal  $p(z)$   
 348 point spread function indicate that Equation 5 is likely too simple, and the  
 349 existence of an efficient computational method to predict  $\tilde{q}_c^{sim}$  for any  $q_c$   
 350 remains an open question.

#### 351 4. Results of the New Algorithm with Standard Misfit

352 A suite of CPT tip resistance ( $q_c$ ) data for known sand-clay layered profiles  
 353 from calibration chamber tests performed by de Lange (2018) at Deltares  
 354 (de Lange, 2018) were used to test the new algorithm with the standard misfit  
 355 function. Details on how the digitized data and reported sample preparation  
 356 were used to estimate  $q_c^{true}$  are provided in Appendix A. Before calculations  
 357 were performed,  $\tilde{q}_c^{meas}$  and  $q_c^{true}$  profiles were converted to normalized cone  
 358 tip resistance ( $q_{c1n}$ ) profiles, where  $q_{c1n}$  is computed as:

$$q_{c1n} = C_N \frac{q_c}{P_a} \quad (6)$$

359 where  $P_a$  is atmospheric pressure in the same units as  $q_c$  and  $C_N$  is a unitless  
 360 overburden correction factor computed per the procedure in (Boulanger and  
 361 Idriss, 2016). Furthermore, data in the upper 0.1 meters of the soil profiles  
 362 was excluded from the analyses presented herein because it contained unin-  
 363 tended artifacts caused by experimental testing limits and was considered to  
 364 be unreliable.

365 It should be noted that all soil models in the de Lange (2018) report  
 366 only have layered zones that contain layers of the same thicknesses, how-  
 367 ever soil models with varying layer thicknesses do not affect the algorithm’s  
 368 performance beyond what is discussed here. For the sake of designing and  
 369 testing the optimization scheme described in Section 3, we replace our  $\tilde{q}_c^{meas}$   
 370 data (which are “naturally blurred”) with profiles that are “computationally  
 371 blurred” (i.e.,  $\tilde{q}_c^{sim}$ ) by applying the convolution in Equation 5 to known soil  
 372 resistance profiles. In the field of inverse problems, this is an example of an  
 373 *inverse crime*, and it is done to test and verify algorithms in a more con-  
 374 trolled setting by removing a source of error from these computational tests.  
 375 This may mean that in practice, the new algorithm is less likely to yield an  
 376 accurate estimate of  $q_c^{true}$ , or that it is more sensitive to the initial  $q_c$  profile  
 377 guess used to begin particle swarm optimization. See the software in Section  
 378 8 and Appendix A for more implementation details.

379 The first soil profile we used to test the new algorithm with the standard  
 380 misfit was Soil Model 4 CPT 2 from the start-up phase of the de Lange (2018)  
 381 data (Section C of the report). This soil profile features 80-mm-thick layers  
 382 of alternating stiffnesses, which are thin enough that existing algorithms  
 383 struggle to correct for thin-layer effects (Yost et al., 2021a). The results are  
 384 shown in Figure 7, comparing the automatically generated initial guess for  $q_c$   
 385 (red solid, left) to the final optimized  $q_c^{inv}$  resulting from the new algorithm  
 386 with the standard misfit function (solid red, right), which is much closer  
 387 to the true resistance profile,  $q_c^{true}$ , based on the known experimental soil  
 388 layering (solid blue). The measured  $\tilde{q}_c^{meas}$  (dashed blue) deviate noticeably  
 389 from the  $\tilde{q}_c^{sim}$  calculated using the initial resistance profile guess (dashed red,  
 390 left), but  $\tilde{q}_c^{sim}$  is extremely close to the computationally simulated  $\tilde{q}_c^{sim}$   
 391 of the final optimized resistance profile  $q_c^{inv}$ . All computations were done in  
 392 serial on an Intel i7 8th generation quad core processor with 8GB DRAM in  
 393 MATLAB 2020a. The algorithm has a fairly good initial guess for  $q_c$  following

394 the procedure in Section 3, but the final best profile,  $q_c^{inv}$ , matches  $q_c^{true}$   
395 extremely well. The only discrepancy is at the very top of the profile where  
396 the  $q_c^{true}$  profile shows a gradient rather than a constant value.

397 Although the new algorithm is only designed to fit piece-wise constant  $q_c$   
398 profiles, it is still able to fit smooth transitions by approximating them with  
399 a stair step-like pattern, the granularity of which depends on the AOI and  
400 LOO parameters passed to the algorithm. Figure 8, which was calculated by  
401 the new algorithm with the standard misfit, shows an example of this. The  
402 default parameters result in an initial  $q_c$  profile guess with several unnecessary  
403 layers at the top of the profile (solid red, left), which were then removed in  
404 the final  $q_c^{inv}$  (solid red, right). This required changing the LOO procedure  
405 to include so-called “absolute thresholding.” This means that, rather than  
406 just removing any layer that does not decrease the misfit much relative to  
407 the current misfit value, the LOO algorithm also removes any layer that only  
408 contributes to the misfit function in a very small region of the profile. Note  
409 that this increased the misfit score from 0.003560 to 0.014702, which can be  
410 interpreted to mean that the final  $q_c^{inv}$  profile has a larger misfit by a roughly  
411 a factor of four compared to the initial  $q_c$  guess, although the final profile is  
412 more physically realistic.

413 The algorithm performed very well on most of the suite of calibration  
414 chamber soil models, but in models with a cluster of very thin layers, the  
415 algorithm skips fitting the last several layers. Figure 9 provides a clear ex-  
416 ample (see also A.13 and A.14). Even though the difference between the  
417 output profile from the new algorithm with the standard misfit,  $q_c^{inv}$ , and  
418  $q_c^{true}$  is large in these cases, the difference between  $\tilde{q}_c^{meas}$  and  $\tilde{q}_c^{sim}(q_c^{inv})$  (in-  
419 dicated by dashed lines) is quite small. The small misfit in this region is  
420 due to the choice of the blurring function, as we can see in the bottom lay-  
421 ers in Figure 1. Considering this limitation, the algorithm performs rather  
422 well. Although this result might be improved by including more layers in this  
423 region in the initial  $q_c$  guess, it is unlikely that all of the layers will be recov-  
424 ered. A blurring function that is more in accord with the physical process  
425 might prevent missed layers in regions with many very thin layers, but this is  
426 not guaranteed. Another approach is to use the new algorithm with the log  
427 misfit function, which emphasizes small differences more than the standard  
428 misfit function when close to the global minimum. More results comparing  
429 the new algorithms with the standard and log misfit are presented in Section  
430 5.

431 To alert software users when it appears there may be a similar scenario

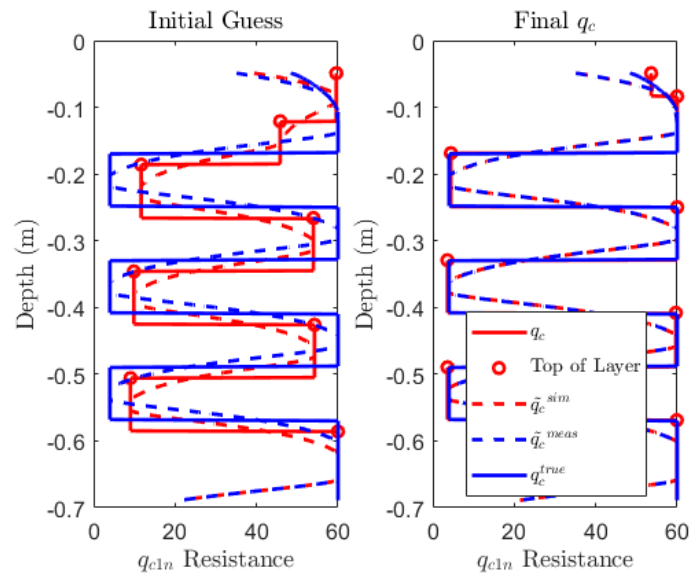


Figure 7: A known CPT tip resistance profile (solid blue), Soil Model 4 CPT 2 from the de Lange (2018) report, section C, compared with a measured  $q_c$  profile (dashed blue), an initial  $q_c$  guess (left, solid red), and the final  $q_c^{inv}$  after coordinate descent, LOO, AOI and particle swarm (right, solid red) show the performance of the method after the code ran for 30 seconds.

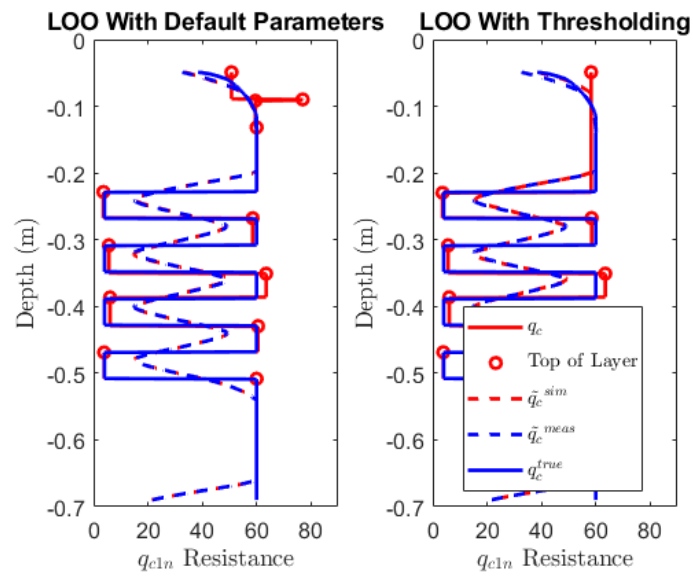


Figure 8: Soil Model 2 CPT 2 from the de Lange (2018) report, Section A. The new algorithm with run with default parameters yielded  $q_c^{inv}$  (left, solid red) that differs from  $q_c^{inv}$  resulting from the new algorithm run with a LOO parameter set to remove layers more aggressively (right, solid red). This example took 80 seconds to run.

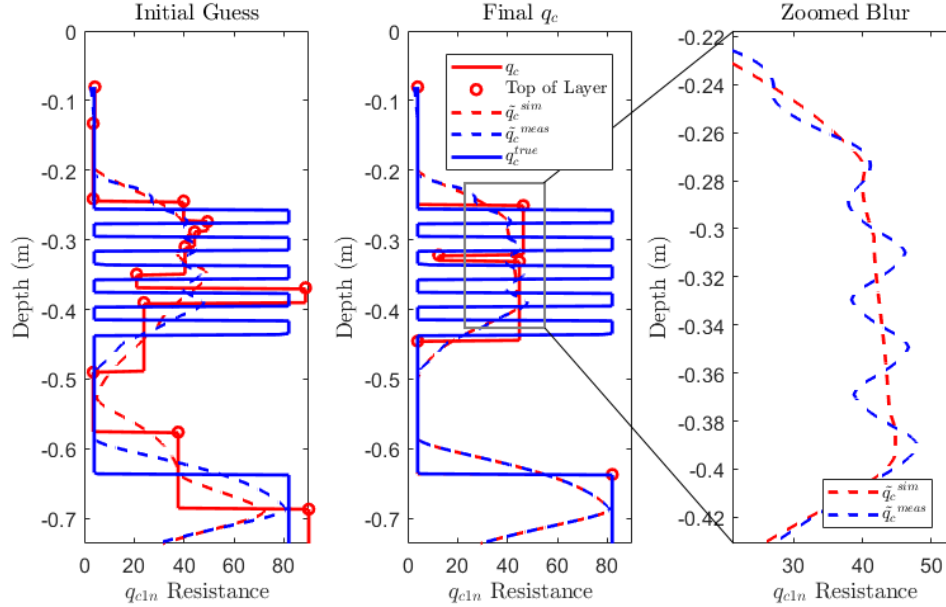


Figure 9: Soil Model 9 CPT 3 from de Lange (2018), Section A. The automatically generated initial  $q_c$  profile guess (left, red solid) and the final updated  $q_c^{inv}$  profile (center, red solid) both miss several layers that are in  $q_c^{true}$  (blue solid). A zoom in of this region (right) shows small-scale differences between  $\tilde{q}_c^{meas}$  and  $\tilde{q}_c^{sim}(q_c^{inv})$ . This process took 40 seconds to run.

432 with many thin layers averaged together and missed, the provided software  
 433 implementation of the new algorithm includes automatically generated warn-  
 434 ings indicating where the algorithm is potentially leaving out a distinct soil  
 435 layer. This is done by dividing up the entire depth profile into regions accord-  
 436 ing to where  $\tilde{q}_c^{sim}(q_c^{inv})$  overestimates or underestimates  $\tilde{q}_c^{meas}$ , and looking  
 437 at the ratio of the misfit function to the signed difference. In practice, this  
 438 seems to work very well even on the example shown in Figure 9. So, it could  
 439 be incorporated into future improvements of the AOI method.

## 440 5. Results of the New Algorithm with Log Misfit

441 We found that the new algorithm with the log misfit function posed in  
 442 Equation 4 was better suited to accurately refine small details of thinly lay-  
 443 ered profiles than the new algorithm with the standard misfit. Figure 10

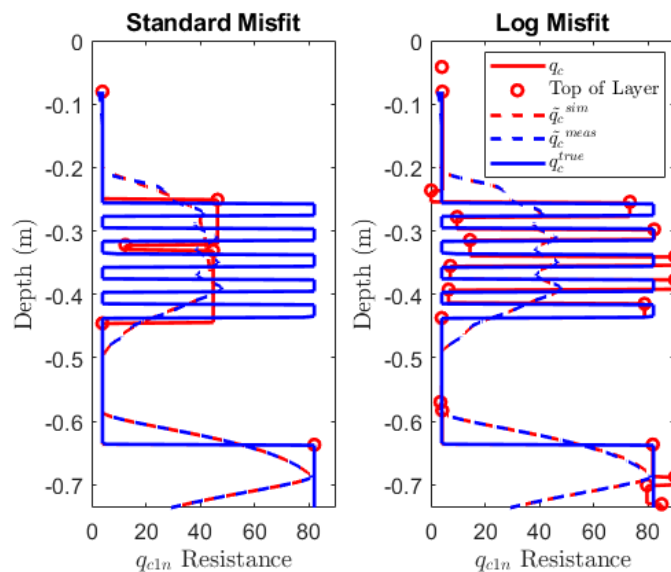


Figure 10: Soil Model 9 CPT 3 from de Lange (2018), Section A. The  $q_c^{inv}$  solution using the standard misfit (solid red, left) is much less accurate than the solution using the log misfit (solid red, right). The standard misfit algorithm took 40 seconds to run, while the log misfit algorithm took approximately 7 minutes.

444 shows the result of this log misfit function applied to the most difficult of  
 445 our previous examples. Note how the  $q_c^{inv}$  resulting from the standard misfit  
 446 (left, solid red) misses many layers in the true resistance profile (solid blue),  
 447 while the  $q_c^{inv}$  resulting from the log misfit (right, solid red) detects every  
 448 single thin layer. For both the standard misfit and the log misfit  $\tilde{q}_c^{sim}(q_c)$  was  
 449 extremely close to  $\tilde{q}_c^{meas}$ , showing that the detailed refinement done by the  
 450 new algorithm with the log misfit was necessary to truly match the thin soil  
 451 profile.

452 Note that minor modifications to LOO and AOI were implemented to  
 453 accommodate the misfit function taking on negative logarithmic values. We  
 454 found it necessary to limit the number of layers that could be added with  
 455 each use of AOI to a small number (three worked well) since marginal im-  
 456 provements to  $\tilde{q}_c^{sim}(q_c)$  can significantly impact the log misfit function. This  
 457 means that, when close to the global minimum, it becomes increasingly diffi-  
 458 cult to add new layers as the algorithm progresses. The new algorithm with  
 459 the log misfit takes between 5 and 10 minutes to run depending on how com-  
 460 plex the actual resistance profile is, as well as how many layers are added

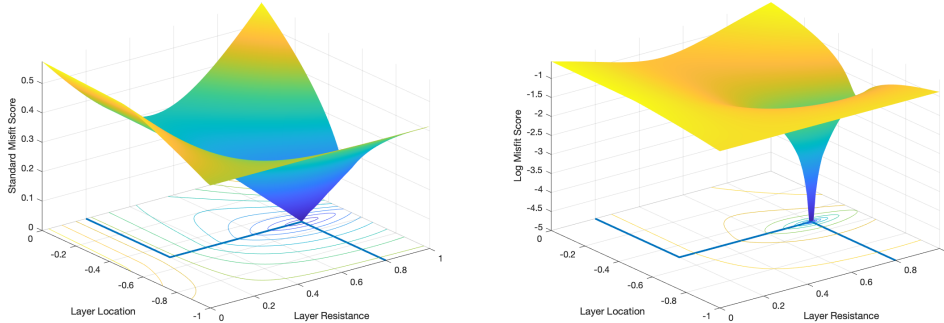


Figure 11: The standard misfit (left) and log misfit (right) are shown for all possible two-layer profiles with the  $q_c^{true}$  value of the top layer known.

461 each iteration. While it is more computationally expensive than the standard  
 462 misfit, the log misfit is also more accurate and robust.

463 Why the log misfit outperforms the accuracy of the standard misfit can  
 464 be understood as follows. Since the minimum of the standard misfit func-  
 465 tion is unique and equals zero, application of a log transform preserves the  
 466 global minimizer (i.e. the best possible  $q_c^{inv}$ ), which becomes the only loca-  
 467 tion where the log misfit approaches negative infinity. This guarantees that  
 468 as we approach the  $q_c^{true}$  profile, PSO should not terminate due to insuffi-  
 469 cient decrease in the objective function value (i.e. the new algorithm with  
 470 the log misfit continues to refine  $q_c$ ). This also has the effect of flattening  
 471 out any local minima that are far from the global minimum, making it even  
 472 easier for PSO to ignore those shallow minima and gravitate towards the  
 473 global minimum. In Figure 11, we compare the contour plots of both mis-  
 474 fit functions for a simple two-layer profile problem. Note that the original  
 475 misfit function has very elongated contours around the minimum. Like the  
 476 well-known Rosenbrock function, this elongated feature can cause numerical  
 477 optimization algorithms to approach the minimum very slowly. Under the  
 478 log transform, the region with elongated contours only has shallow decreases  
 479 in the misfit function with more circular contours closer to the minimum.

## 480 6. Extensions of Method

481 Additional improvements in the computational efficiency of this method  
 482 could be achieved through the use of parallel computing and GPU comput-



483 ing. While parallel computing in science and engineering has historically  
484 focused on massive problems running on large computer clusters, multi-core  
485 architectures and rapidly improving graphics cards are now widely available  
486 in laptops and desktops accessible to most engineers.

487 Further improvement in the accuracy of the method could be achieved  
488 through a more realistic representation of the measurement blurring process  
489 (i.e., the influence of thin-layer effects on  $\tilde{q}_c^{meas}$ ). For our examples, we use  
490 a simple convolution with a smooth, asymmetric pointspread function (that  
491 is,  $\tilde{q}_c^{meas}$  at a given depth is affected more by  $q_c^{true}$  below the depth than  $q_c^{true}$   
492 above the depth). However, calibration chamber test data and high-fidelity  
493 material point method numerical simulations of soil displacement during pen-  
494 etrometer testing (Zambrano-Cruzatty and Yerro, 2020; Yost et al., 2021b)  
495 reveal much more complex physics, suggesting that a simple convolution with  
496 a single pointspread function is inadequate in some scenarios.

497 Moving forward, we aim to develop a method based on high-fidelity simu-  
498 lations to yield a more physically realistic computational model of this blur-  
499 ring process that is computationally fast to apply. For example, a neural  
500 network trained to mimic the numerical blurring of any  $q_c^{true}$  profile will also  
501 be computationally cheap to evaluate. Speed is an important feature in the  
502 layer optimization algorithm, since we expect to calculate  $\tilde{q}_c^{sim}$  many times  
503 within each iteration. The primary drawback is the large amount of training  
504 data required, which is experimentally challenging to acquire and computa-  
505 tionally taxing to generate via high-fidelity simulation.

## 506 7. Discussion and Conclusions

507 Thin-layer correction for CPT data can be posed as an inverse problem,  
508 similar to other signal deblurring problems. Our tests indicate that solving  
509 for a  $q_c$  profile mathematically represented by an independent tip resistance  
510 value at every depth (as done in Boulanger and DeJong (2018)) does not  
511 reliably yield improved data quality (Yost et al., 2021a), even when one adds  
512 regularization to enhance blocky layers of stratigraphy (a common strategy  
513 in image deblurring). We pose this inverse problem in a new way, searching  
514 for a finite number of subsurface layers, each having a thickness and uniform  
515 resistance that must be found such that  $\tilde{q}_c^{sim}$  most closely matches  $\tilde{q}_c^{meas}$ .  
516 Our tests indicate this new formulation of the inverse problem is better able  
517 to identify thin interbedded layers in the soil profile compared to previous  
518 methods tested in Yost et al. (2021a).

519 We developed open-source software that takes as inputs a measured CPT  
520 tip resistance profile ( $\tilde{q}_c^{meas}$ ) together with code to mimic the natural “blurring”  
521 of the true  $q_c$  profile due to the stress bulb that forms around the cone  
522 penetrometer tip. It estimates a piecewise constant  $q_c$  profile that is expected  
523 to result in data similar to the measured data, up to a desired tolerance.  
524 With appropriate settings, this method will correct for thin-layer effects dur-  
525 ing CPT soundings, although there may be other sources of measurement  
526 error that remain uncorrected. Our tests indicate that typical profiles char-  
527 acterized by a moderate number of depths at which  $q_c^{meas}$  is recorded (i.e.,  
528 with several hundred values) can be corrected using the new algorithm with  
529 the standard misfit within 1-2 minutes on a standard laptop with one core.

530 This software has limitations. In many cases, the resulting  $q_c^{inv}$  profile will  
531 be simpler and smoother than what we might expect from the true resistance  
532 profile if we were to have other soil profile characterization data (e.g. core  
533 samples). However this tendency towards simplified models can easily be ad-  
534 justed by the user’s settings. The existence of a fast computational procedure  
535 that closely mimics this “natural blurring” of the true soil resistance profile  
536 remains an open question, so we use a simple convolution (5) to compute the  
537 “blurred” data,  $\tilde{q}_c^{sim}$ . Future research will focus on quantifying uncertainty  
538 in the corrected  $q_c^{inv}$  profiles, and improving the blurring functions that are  
539 used to represent a wider range of geotechnical scenarios.

540 For particularly complex soil stratigraphies, we suggest users take ad-  
541 vantage of the new algorithm with the log misfit, which balances regions  
542 with large-scale and fine-scale features contributing to the misfit. Our tests  
543 reveal this yields more accurate and robust profiles, better reflecting subsur-  
544 face stratigraphy, with more thin layers correctly identified. Our open-source  
545 software includes both the standard and log misfits. Note that the new algo-  
546 rithm with the log misfit requires a longer run time, typically 5-10 minutes  
547 for CPT soundings with several hundred depth points. By providing soft-  
548 ware for both formulations, users can decide which version to apply based on  
549 tradeoffs in computing time, accuracy, and assumed complexity of the soil  
550 stratigraphy.

## 551 8. Data Statement

552 Multiple examples in this paper were performed using data that is avail-  
553 able through a technical report from Deltares (de Lange, 2018). Code for

554 every algorithm and example in this paper are publicly available at [https:](https://github.com/jonc7/Soil-Layer-Optimization)  
555 [//github.com/jonc7/Soil-Layer-Optimization](https://github.com/jonc7/Soil-Layer-Optimization).

## 556 9. Acknowledgements

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561 sions or recommendations expressed in this material are those of the authors  
562 and do not necessarily reflect the views of the NSF.

## 563 Appendix A. Additional Soil Model Information

564 In the de Lange (2018) study, several different soil models consisting of  
565 layered sand-clay profiles were constructed and CPTs were performed in these  
566 models at various in-situ stress conditions. Sand layers in the soil models were  
567 prepared with target relative densities of either 30% (loose) or 60% (dense).  
568 Uniform or “reference” sand models were also prepared with the same target  
569 relative densities. CPTs performed in the reference sand models therefore  
570 provided a good estimate of the “true” resistance profile  $q_c$  for the sand  
571 layers in the corresponding layered soil models. Note that due to variation  
572 in experimental preparation of the soil models, the sand relative densities in  
573 the layered models did not always match well with that of the reference sand  
574 model. For the purposes of this study, only the layered soil models that were  
575 relatively good matches were considered [this included several layered models  
576 presented in the “Test Results” section (Section A) and Soil Model 4 from  
577 the “Start-Up Phase” section (Section C) of the de Lange (2018) report].  
578 Furthermore, no uniform/reference clay soil models were considered by de  
579 Lange (2018), therefore, the “true”  $q_c$  in the clay was estimated based on  
580 the minimum  $q_c$  observed in the thickest of the clay layers in the de Lange  
581 (2018) experiments. With the estimated “true”  $q_c$  for both the clay and  
582 sand layers, and knowledge of the layer depths and thicknesses,  $q_c^{true}$  profiles  
583 for the layered soil models were constructed for comparison to the measured  
584 CPT tip resistance profiles,  $\tilde{q}_c^{meas}$ , and the corrected tip resistance profiles  
585 calculated by the new algorithms,  $q_c^{inv}$ .

586 Figures A.12, A.13 and A.14 show results of the simple procedure to  
587 pick an initial profile guess, the final  $q_c^{inv}$  for the standard objective function,

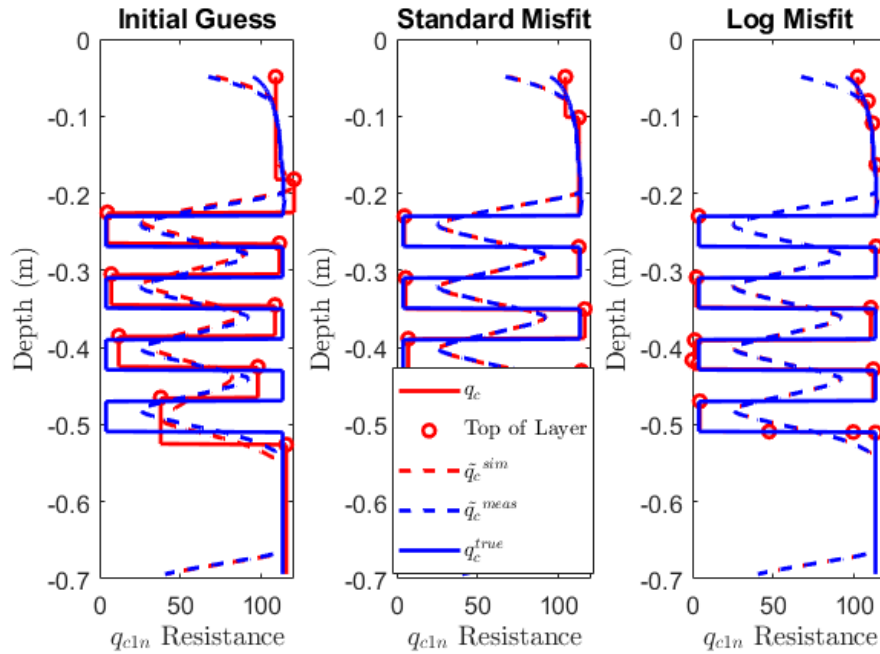


Figure A.12: Soil Model 4 CPT 3 from the de Lange (2018) report, Section A. The automatically generated initial  $q_c$  profile (left, solid red) is compared to the final  $q_c^{inv}$  profiles resulting from the new algorithm with standard misfit (center, solid red) and the log misfit (right, solid red). The algorithm with the standard and log misfits took 50 seconds and over 7 minutes, respectively, to run.

588 and the final  $q_c^{inv}$  for the log objective function for three laboratory datasets  
 589 from the de Lange (2018) report.

## 590 References

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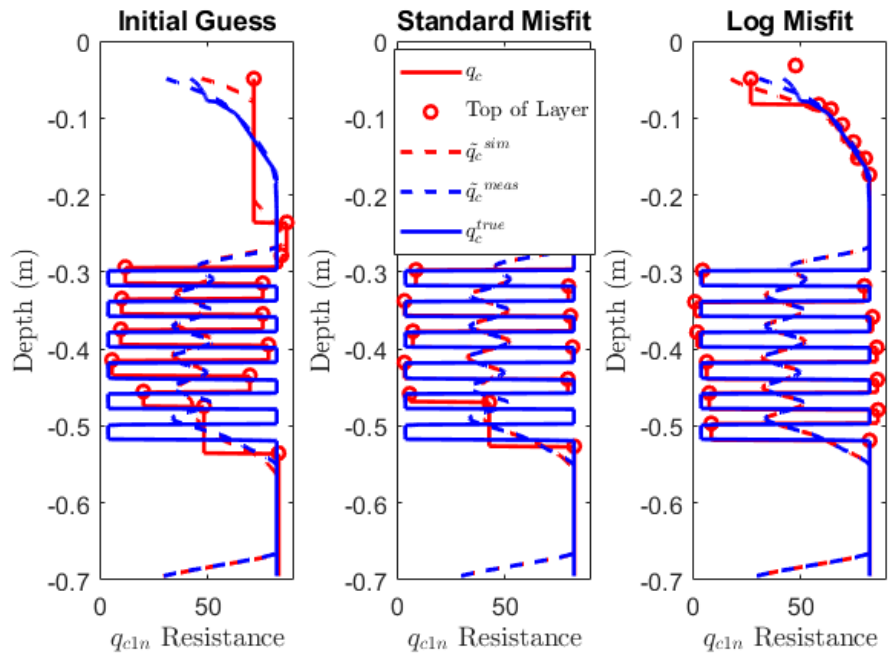


Figure A.13: Soil Model 3 CPT 3 from the de Lange (2018) report, Section A. The automatically generated initial profile (left, solid red), the final  $q_c^{inv}$  profiles from the new algorithm with the standard misfit (center, solid red) and the log misfit (right, solid red) are shown. The new algorithm with the standard misfit took 100 seconds, and 10 minutes with the log misfit.

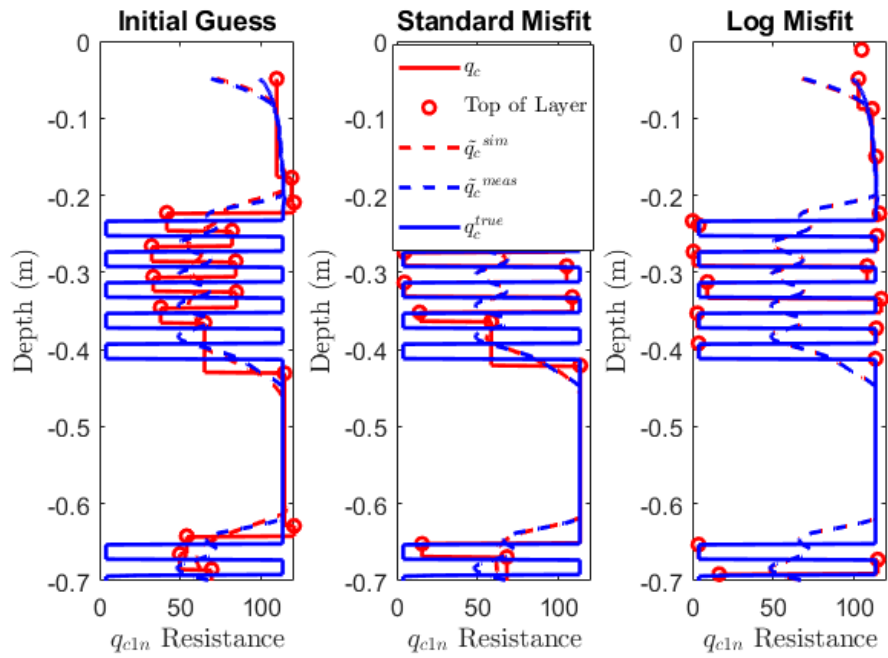


Figure A.14: Soil Model 8 CPT 3 from the de Lange (2018) report, Section A. The automatically generated initial  $q_c$  profile (left, solid red), the final  $q_c^{inv}$  profiles generated by the new algorithm with the standard misfit (center, solid red) and log misfit (right, solid red) are shown. The algorithm with the standard misfit took 70 seconds, and the log misfit took over 11 minutes.

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